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Summary. Plastic deformation can occur by a collective motion of atoms as crystal dislocations. Evidence for dislocation creep in Earth is abundant although other mechanisms such as diffusional creep (see Chapter 8) dominate under some conditions. The rate of deformation due to dislocation motion is proportional to dislocation density and velocity (the Orowan equation).

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The deformation recorded in the rocks occurs when the rocks break (fail) or flows viscously (a type of crystal-scale failure) or plastically (also a crystal-scale failure process). Elastic deformation is the dominant form of deformation at shallow depths in the crust and lithosphere because both the temperature and pressure are low.

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Deformation of a polycrystalline material is associated with grain boundary sliding. Large-strain plastic flow involving grain-boundary

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sliding is sometimes referred to as superplastic flow. Materials science models of superplastic flow are reviewed and some geological significance is discussed.

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